

A STUDY OF EARTHQUAKE MAGNITUDES AND THEIR  
RELATION TO THE LAW OF SEISMIC SPECTRUM SCALING

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## DIGEST

It is known that short-period body-wave magnitude,  $m_b$ , is not a reliable indicator of the strength of very large earthquakes. For these earthquakes the surface-wave magnitude,  $M_S$ , appears to be a more reliable index of the strength. A possible solution to the body-wave magnitude problem would be to determine the magnitude from the lower-frequency part of the P-wave spectrum, such as would be obtained from the amplitude of P as recorded by long-period seismographs. This P-wave magnitude is called  $m_B$ .

Predicted  $m_b-M_S$ ,  $m_B-M_S$  and  $m_b-m_B$  curves were constructed from spectra obtained from some of the source models in current vogue; (Aki's  $\omega^{-2}$ , A and B Models; Archambeau's  $V_r/V_p = \frac{1}{2}$ ,  $R_s/R_o = 2$  Model; Berckhemer and Jacob's Model; and Brune's Partial Stress Drop Models  $e = 0.1$  and  $0.01$ ). These models all predict a slope near 1 for the  $m_b$  versus  $M_S$  and  $m_B$  versus  $M_S$  curves at large magnitude. However, the different models predict widely varying values for the absolute difference between  $m_B$  and  $m_b$ ; this difference is a function of the slope of the spectra at intermediate frequencies.

Magnitude data obtained using the first few cycles of P motion did not support the prediction of the models at high

magnitude, that is both  $m_b-M_S$  and  $m_B-M_S$  plots of observed data flatten out at the high magnitude end, although this is not predicted by the mathematical models. However, as the seismograms were analyzed, it was noted that a high percentage of the larger earthquakes had later P phases which were consistently larger than the first arrival at all distances and azimuths. When these "multiple" earthquakes were identified, it was noted that they separated out (smaller  $m_b$  and  $m_B$  values) from the remaining shocks on the  $m_b-M_S$  and  $m_B-M_S$  plots and that the remaining earthquakes supported the predicted curves. Body-wave magnitudes,  $m_b$  and  $m_B$ , were recomputed from the amplitude and period of the later "main event" for "multiple" earthquakes. On  $m_b-M_S$  and  $m_B-M_S$  plots the "simple" and "multiple" earthquakes became compatible, the scatter in the plots was reduced, and the plots supported the predicted curves.

The observed absolute difference between  $m_b$  and  $m_B$  was found to be independent of the nature of the earthquake ("simple" or "multiple") and supported predicted results of only Aki's Model B and Brune's Partial Stress Drop models.

P-wave spectra were computed for earthquakes over an  $M_S$  range of 4.9 to 8.4. These spectra were characterized by two corner frequencies, a slope of -3 for the migration of the

low frequency corner with increasing displacement amplitude, a slope of -3 for the migration of the high frequency corner below  $M_S = 7.1$ , and a vertical migration of the high frequency corner above  $M_S = 7.1$ .

Predicted  $m_b - M_S$ ,  $m_B - M_S$  and  $m_b - m_B$  curves were constructed from the observed spectra. The  $m_b - M_S$  and  $m_B - M_S$  curves obtained from the observed spectra do not support the Hanks and Wyss (1972) assumption that the S-wave spectrum is shifted by an amount  $\alpha/\beta$  to lower frequency with respect to the P-wave spectrum. The data support Savage's (1972) prediction that the spectra should be shifted the other way.

Conventional  $m_b$  is considered to be a questionable parameter for all shallow earthquakes of  $M_S$  greater than  $5\frac{1}{2}$ . Conventional  $m_b$  has little relation to the strength of earthquakes of  $M_S$  greater than  $6\frac{1}{2}$ . Conventional  $m_B$  fares somewhat better in maintaining a relation to the size of the earthquake (at least  $m_B$  increases when  $M_S$  increases) for about another order of magnitude.