Cold-Season Coupled Upper-Level Jet Streaks in the Northeastern United States

Part 1: Weak Dynamic Cases





²Cooperative Institute for Precipitation Systems, Department of Earth and Atmospheric Sciences, Saint Louis University ³NOAA/NWSFO Buffalo, NY



Introduction

- The interaction of transverse circulations associated with two separate upper-level (UL) jet streaks, along with its effects on sensible weather, have been documented in several studies:
 - East Coast cyclogenesis (e.g., Uccellini and Kocin 1987)
- Heavy/banded precipitation (e.g., Belville and Stewart 1983; Junker et al. 1990; Hakim and Uccellini 1992; Funk and **Moore 1995; Melde 1996)**
- Organized severe thunderstorm complexes (e.g., Hamilton et al. 1998; Ashley et al. 2000; Jamski et al. 2000)

SAINT LOUIS

UNIVERSITY

- The term 'coupled jet streaks' refers to the presence of two separate jet streaks juxtaposed in such fashion that the ascending branches of the transverse circulations are collocated with one another, resulting in an enhanced area of upward vertical motion (e.g., Uccellini and Kocin
- This study will investigate coupled UL jet streak occurrences during the cool season (1 October to 31 March) in the northeastern U.S. over 10 seasons (1993 – 2003).



Preliminary examination of the UL flow regime using the North American Regional Reanalysis (NARR) dataset revealed 79 possible coupled jet streak occurrences during the period.

- Using the General Meteorological Package (GEMPAK) with the NARR dataset, plan-view and cross-sectional analyses of the possible occurrences were analyzed to ensure the interaction of the jet streak circulations.
- This revealed 39 coupled jet streak cases, which were then subdivided into weak dynamic (n=20) and strong dynamic (n=19) scenarios.
- The weak dynamic cases (covered in this presentation) were characterized by modest surface circulations (MSLP > 1000 hPa) and open mid-tropospheric

8 March 2002

Regions of 8+ inches of snow

80+ reports of thundersnow

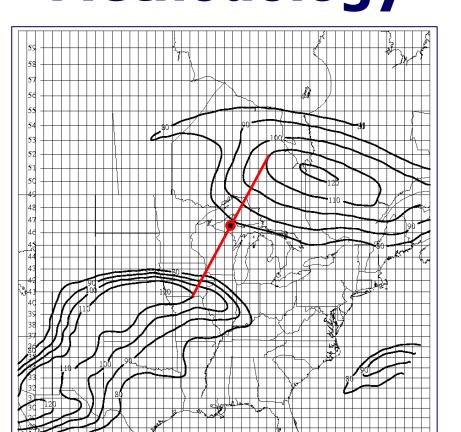
Significant inverted trough

Lake influence unlikely

development at surface (w/c

significant cyclone in vicinity)

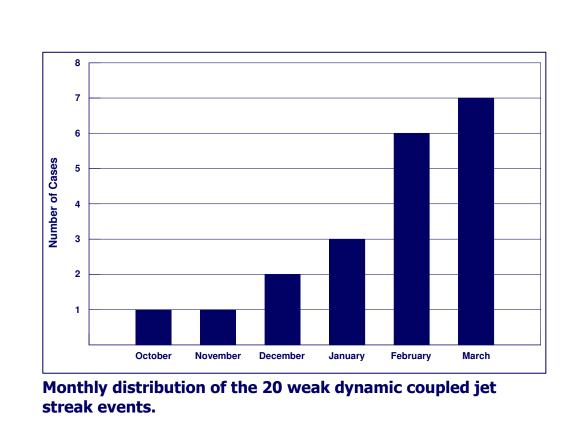
Methodology

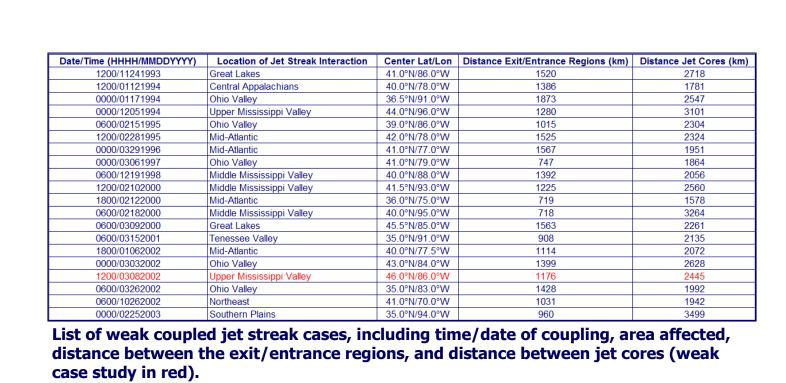


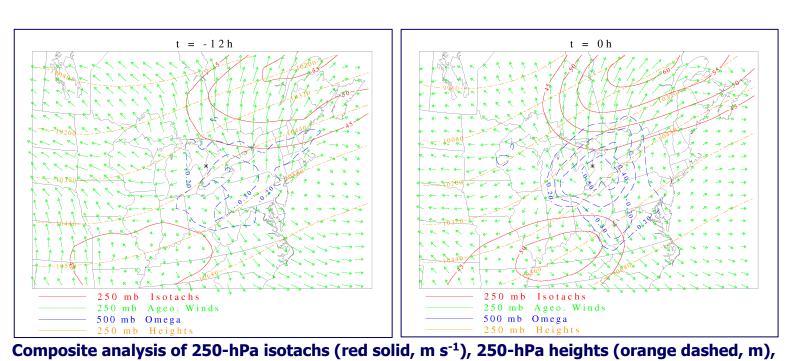
8 Mar 2002, based upon cooperative station data from NCDC

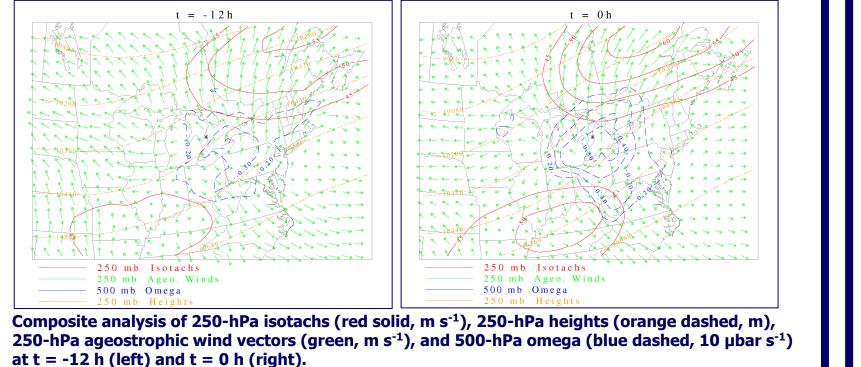
- Center points between the jet streaks were then qualitatively determined by finding the midpoint on a line between the strongest common isotach for the initial coupling time, along with the prior 6- and 12-h time periods (see left).
- The NARR data was objectively analyzed to a 31 x 23 grid with 128 km between gridpoints using the Barnes (1973) objective analysis scheme.
- A 27 x 19 grid was then extracted using the center point between the jet
- Finally, the data was then averaged over the 20 weak cases using the SLUBREW compositing software developed by Moore et al. (1996).

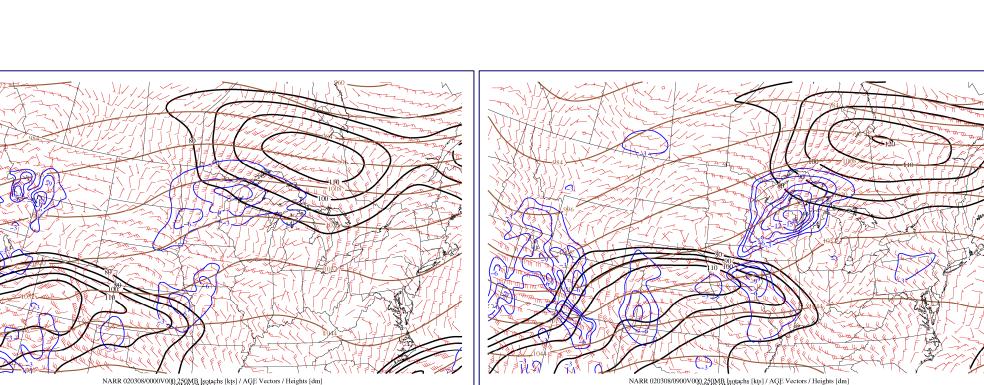
Composites

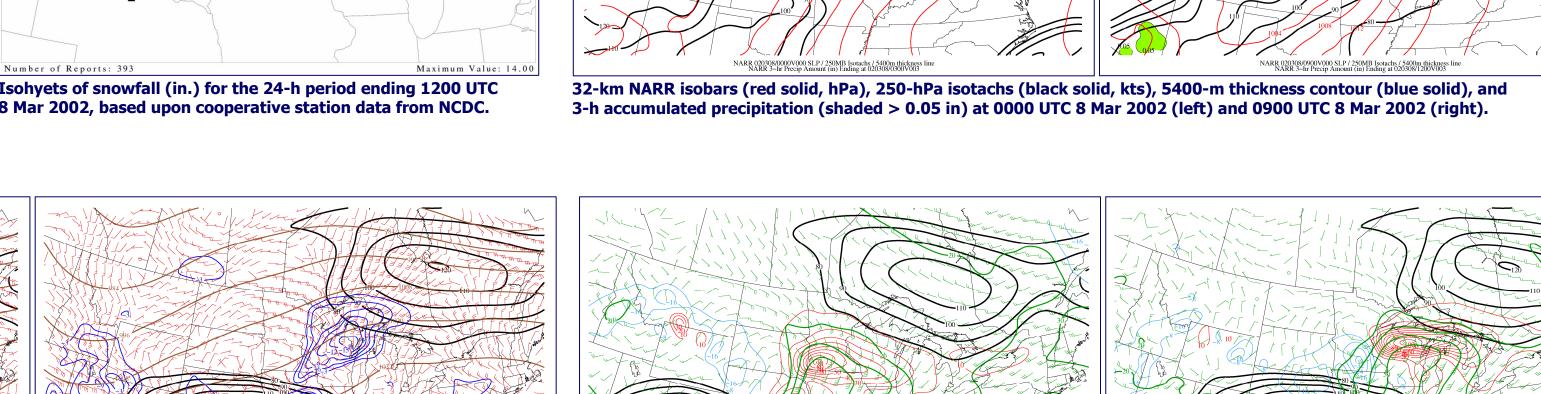








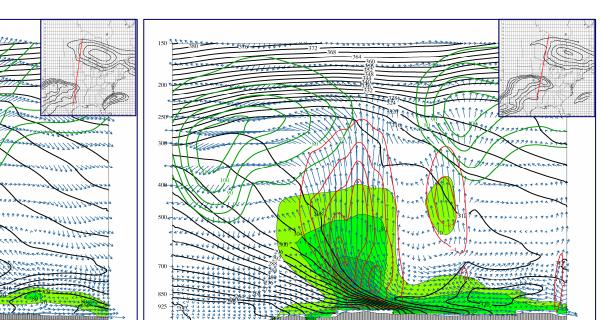




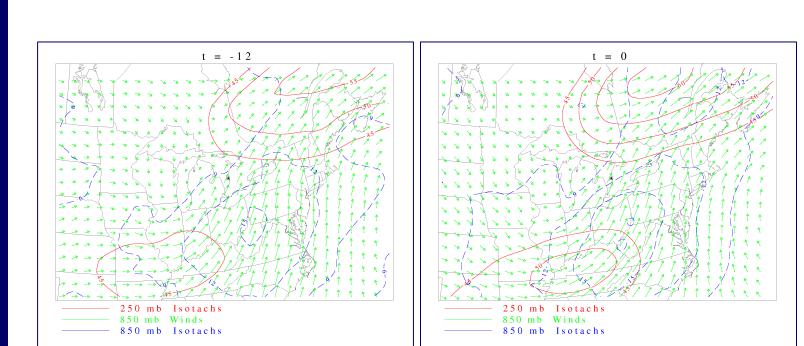
32-km NARR 250-hPa heights (brown solid, dkm), 250-hPa isotachs (black solid, kts), 250-hPa ageostrophic wind barbs (red), and 500-hPa omega (blue solid, μ bar s⁻¹) at 0000 UTC 8 Mar 2002 (left) and 0900 UTC 8 Mar 2002 (right).

32-km NARR 250-hPa isotachs (black solid, kts), 850-hPa isotachs (green solid, kts), 850-hPa wind vectors (green, kts), and

850-hPa θ_e advection (red solid [+10⁻¹ K hr⁻¹], blue solid [-10⁻¹ K hr⁻¹]) at 0000 UTC 8 Mar 2002 (left) and 0900 UTC 8 Mar 2002



NARR isotachs (green solid, kts), omega (red solid, μ bar s⁻¹), equivalent potential temperature (black solid, K), ageostrophic vertical circulation (blue arrows), and RH \geq 70% (green shading) at 0000 UTC 8 Mar 2002 (left) and 0900 UTC 8 Mar 2002 (right). Inset figures provide the orientation of the cross-section with respect to the



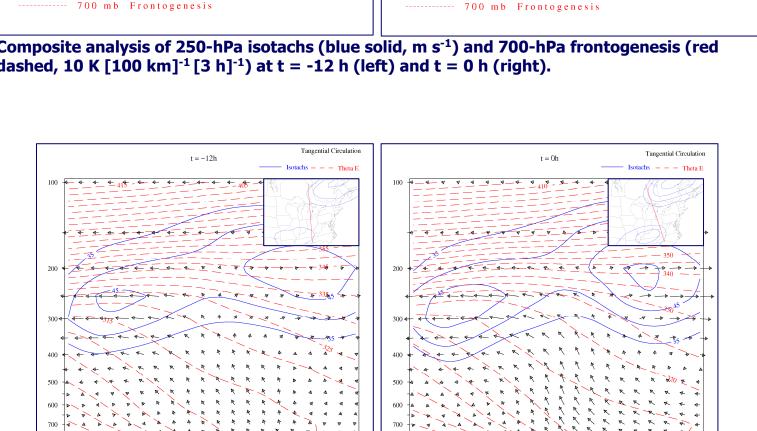
Composite analysis of 250-hPa isotachs (red solid, m s⁻¹), 850-hPa isotachs (blue dashed, m s⁻¹), and 850-hPa wind vectors (green) at t = -12 h (left) and t = 0 h (right).

Composite analysis of 250-hPa isotachs (red solid, m s⁻¹), 850-hPa wind vectors (green), and θ_e advection (blue dashed, K hr⁻¹) at t = -12 h (left) and t = 0 h (right).

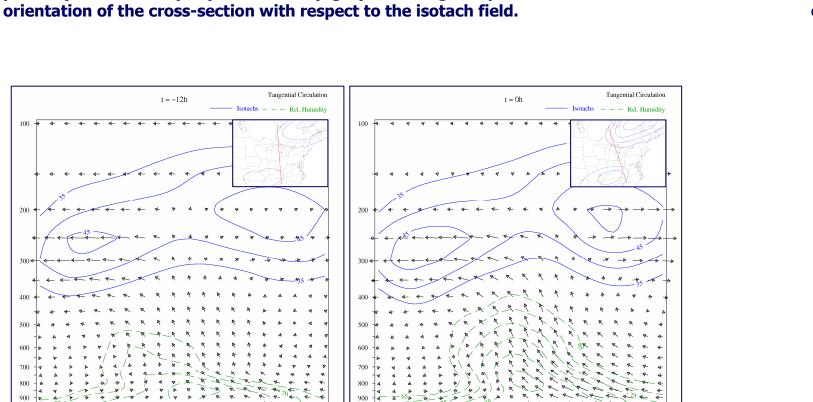
Cross section through the composite coupled jet streaks showing isotachs (blue

μbar s^{-1}) at t = -12 h (left) and t = 0 h (right). Inset figures provide the

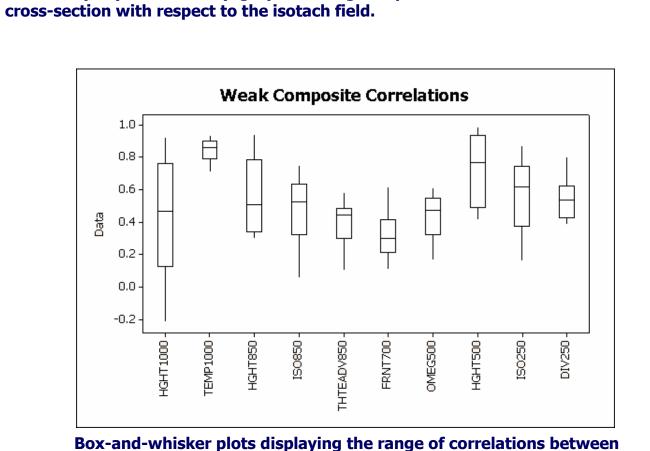
solid, m s⁻¹), ageostrophic circulation (black arrows), and omega (red dashed, 10



Cross section through the composite coupled jet streaks showing isotachs (blue solid, m s⁻¹), ageostrophic circulation (black arrows), and $\theta_{\rm e}$ (red dashed, K) at t = -12 h (left) and t = 0 h (right). Inset figures provide the orientation of the



Cross section through the composite coupled jet streaks showing isotachs (blue solid, m s⁻¹), ageostrophic circulation (black arrows), and relative humidity $\geq 60\%$ (green dashed, %) at t = -12 h (left) and t = 0 h (right). Inset figures provide the orientation of the cross-section with respect to the isotach field.



the composite fields and the individual weak cases. In each case the horizontal line within each box represents the 50th percentile, the top (bottom) line of the box is the 75th (25th) percentile, and the top (bottom) line outside the box represents the 90th (10th) percentile.

Conclusions

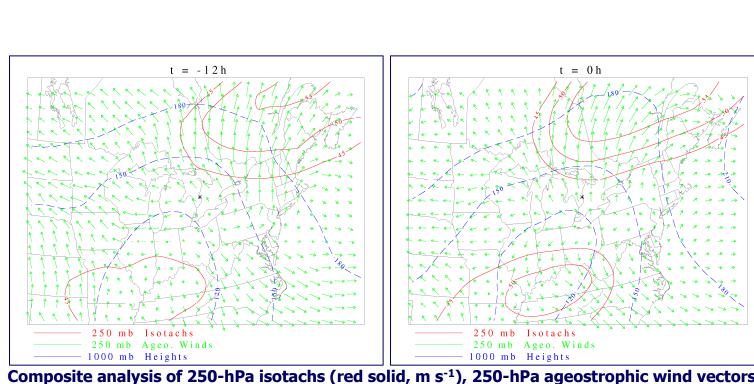
Weak Dynamic Case Study

Number of weak dynamic scenarios increases steadily from winter to spring 1227 (2351) km average distance between jets' entrance and exit regions (jet cores) at time of coupling

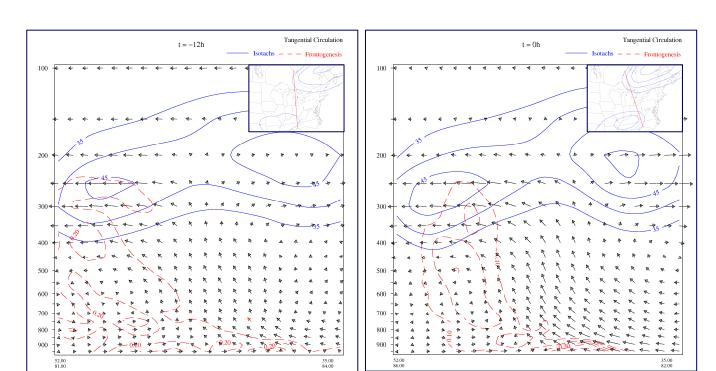
32-km NARR 250-hPa isotachs (black solid, kts), and 850-hPa frontogenesis (red solid, K [100 km]⁻¹ [3 h]⁻¹) at 0000 UTC 8 Mar

- Southern jet 'parallel' to northern jet
- Northern jet retrogrades slightly and strengthens during coupling period, while southern jet progresses slowly eastward and strengthens
- 250-hPa heights fall significantly upwind of jet streaks, but trough remains
- 250-hPa ageostrophic cross-contour flow strengthens during coupling period, resulting in increased/focused upper-level divergence and mid-tropospheric **UVM** in coupling region
- 850-hPa low-level jet and θ_e advection become stronger and better organized in coupling region over coupling period
- 700-hPa frontogenesis region elongates and strengthens underneath entrance region of northern jet during coupling period

- 'Surface' low remains weakly defined during coupling period, while 1000-hPa inverted trough intensifies in vicinity of northern jet's entrance region
- Poleward/upward transport of warm/moist (high θ_e) air (i.e. ageostrophic circulation) becomes better defined over coupling period
- UVM region narrows/strengthens in between jets over coupling period
- Low-level front strengthens over coupling period
- Mid/upper-tropospheric frontogenetical circulation works in concert with coupled jet circulation as frontogenesis extends upward over coupling period
- Moist layer between jets deepens and narrows over coupling period
- Correlations of basic parameters generally better than those of derived parameters, but minimal differences between weak and strong episodes indicate similar properties for both types of interaction
- This project would not have been possible without COMET Partners Project Award S05-52248 and the help of Charles E. Graves (Saint Louis University) who aided the authors in using the SLUBREW compositing software.



(green), and 1000-hPa heights (blue dashed, m) at t = -12 h (left) and t = 0 h (right).



Cross section through the composite coupled jet streaks showing isotachs (blue solid, m s⁻¹), ageostrophic circulation (black arrows), and frontogenesis (red dashed, 10 K [100 km]⁻¹ [3 h]⁻¹) at t = -12 h (left) and t = 0 h (right). Inset figures provide the orientation of the cross-section with respect to the isotach